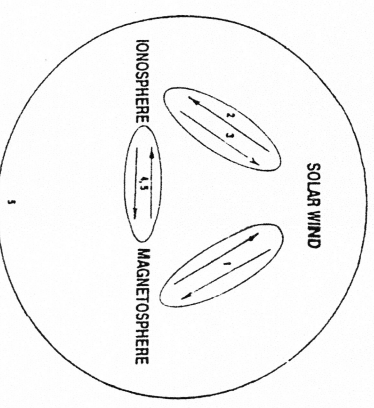


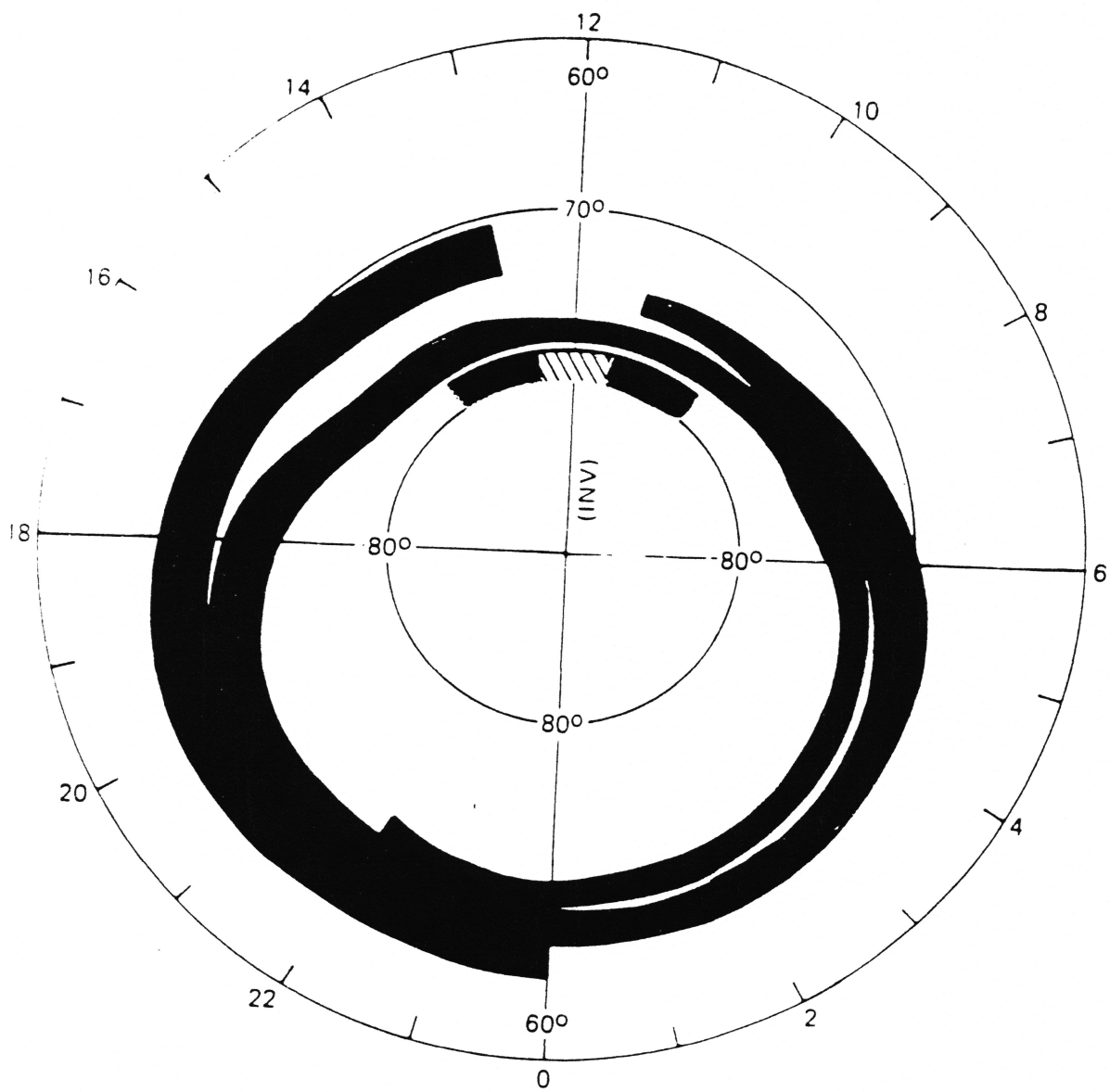
SOLAR WIND - MAGNETOSPHERE - IONOSPHERE COUPLING

LECTURE 4

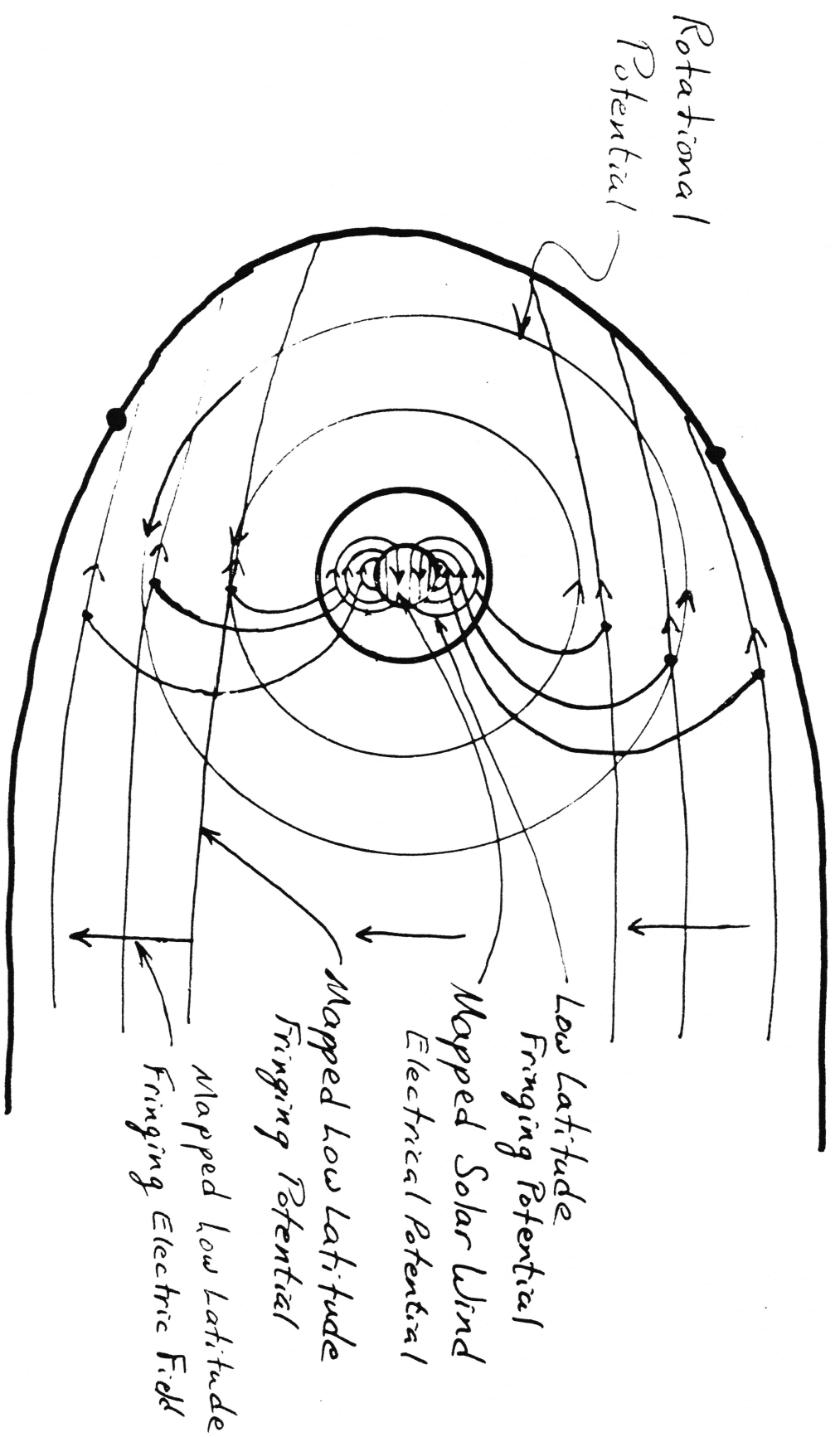
MAGNETOSPHERE - IONOSPHERE COUPLING

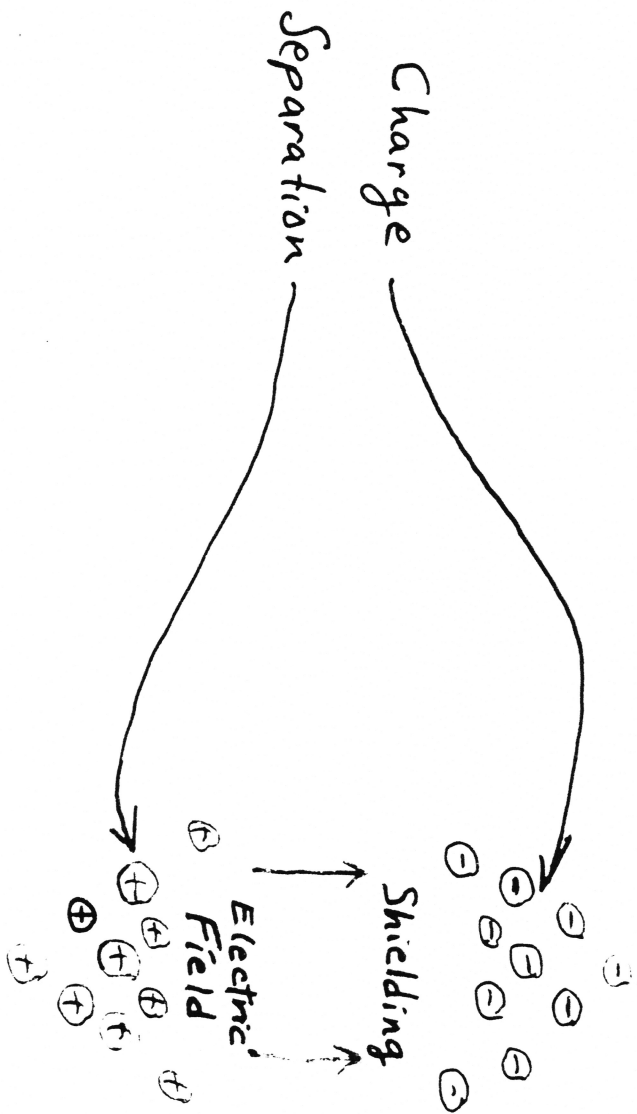


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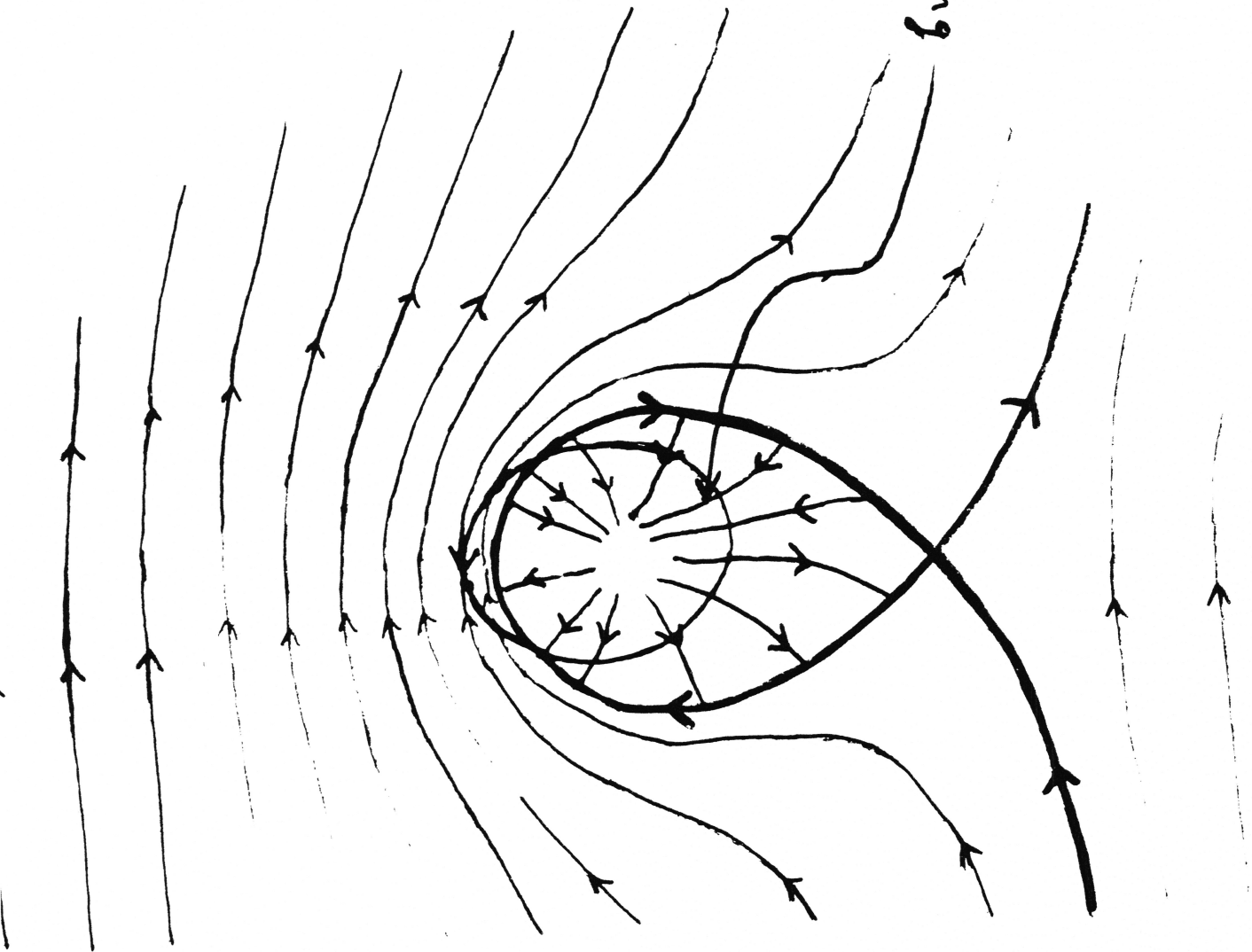


Currents into Ionosphere
Currents Away from Ionosphere



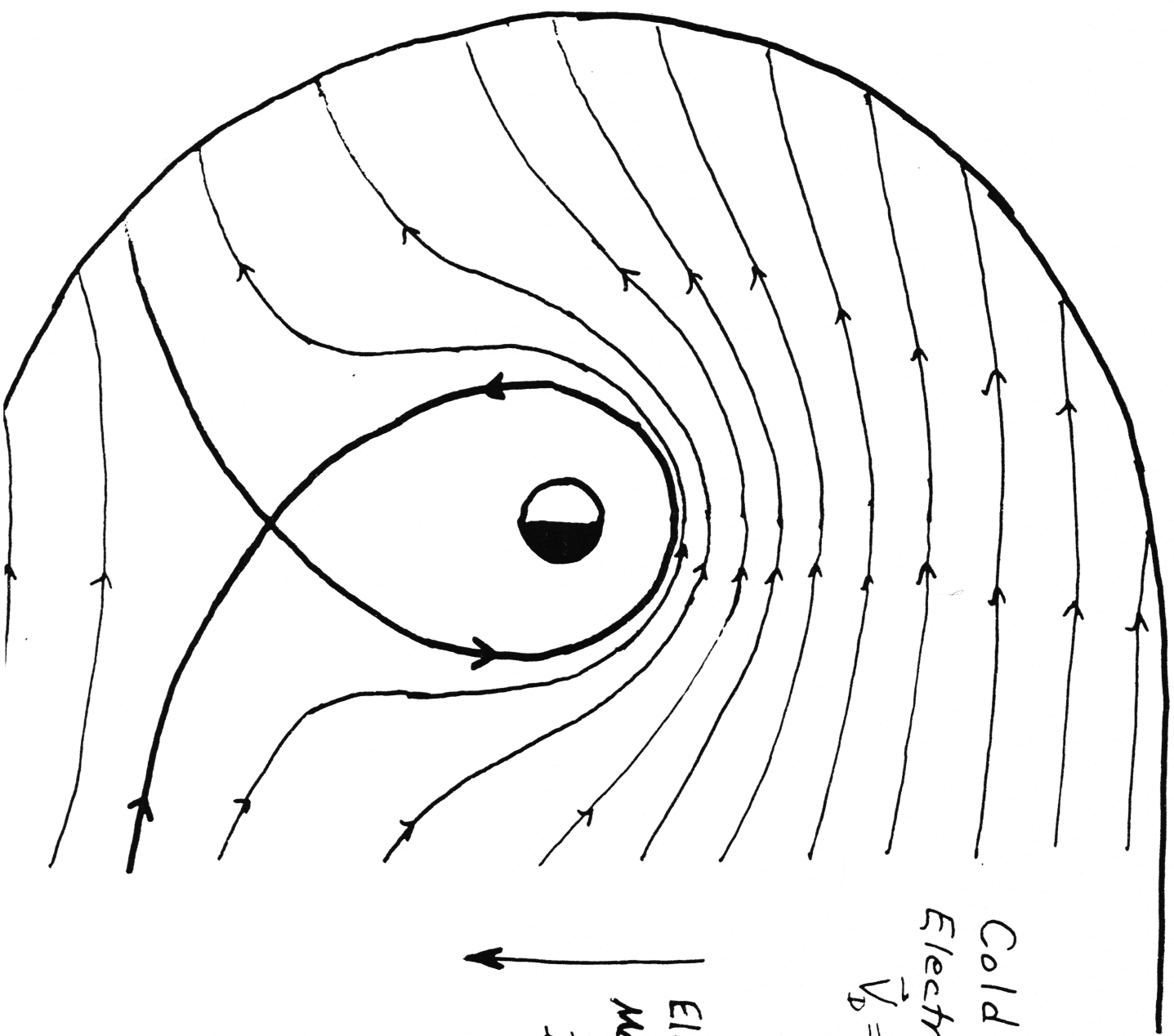


Discharging
Currents



Hot
Ions

$$\vec{V}_p = \frac{\vec{E} \times \vec{B}}{B^2} + \vec{V}_a + \vec{V}_c$$



Cold

Electrons

$$\vec{V}_D = \frac{\vec{E} \times \vec{B}}{B^2}$$

Electric Field
Mapped From
Ionosphere

VASYLIUNAS MAGNETOSPHERE - IONOSPHERE COUPLING EQUATION

Magnetospheric Current

$$\vec{I}_m = eN(\vec{V}_B + \vec{V}_D) - eN\vec{V}_B + \vec{I}_M$$

$$\vec{V}_B = \vec{E} \times \vec{B}_1 / B_1^2$$

Current into Ionosphere from Magnetosphere

$$J_{\parallel} \sin \chi = -\nabla \cdot \vec{I}_m$$

Continuity Equation for Protons

$$\frac{\partial N}{\partial t} + \nabla \cdot N(\vec{V}_B + \vec{V}_D) = 0$$

Combining above gives

$$J_{\parallel} \sin \chi = \nabla \cdot eN\vec{V}_B + e \frac{\partial N}{\partial t}$$

Which can be written as

$$J_{\parallel} \sin \chi = -\nabla \cdot \sum_H^* \frac{\vec{B}_1}{B_1} \times \vec{E} + e \frac{\partial N}{\partial t}$$

$$\sum_H^* = \frac{eN}{B_1}$$

Ionospheric current continuity equation

$$\nabla \cdot (\sum \cdot \vec{E}) = J_{\parallel} \sin \chi$$

Combining above gives

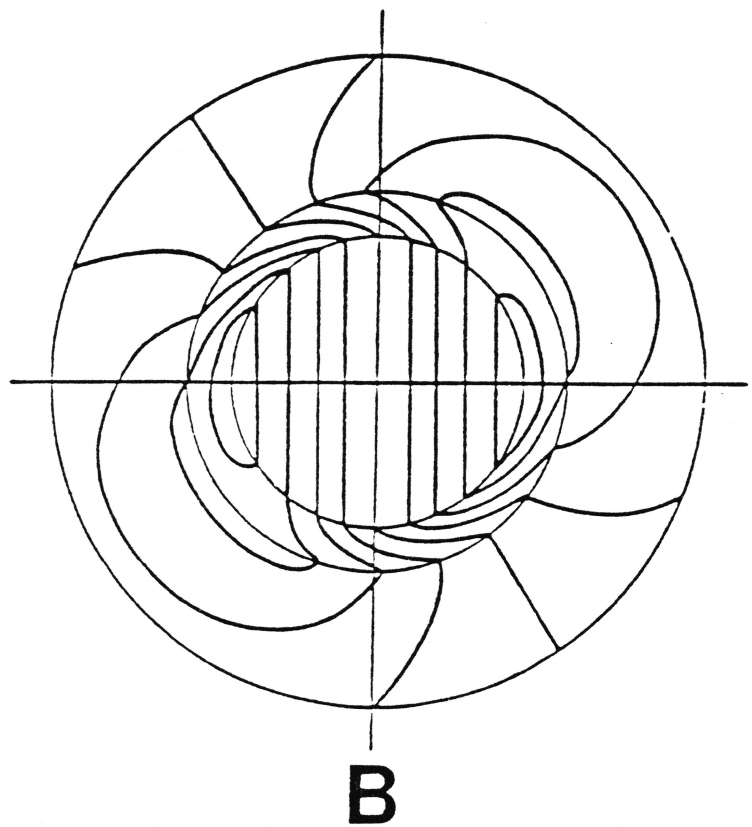
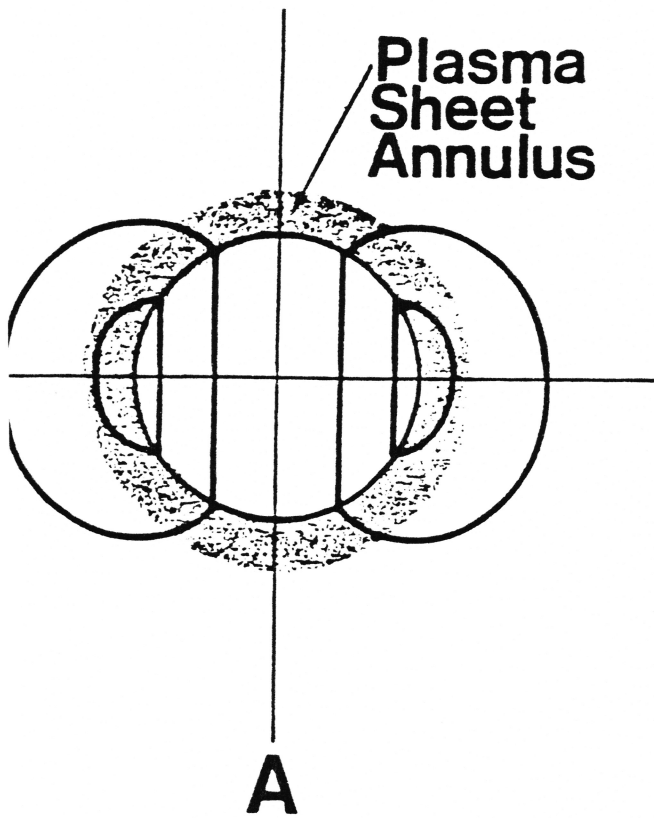
$$\nabla \cdot (\sum + \sum_H^*) \cdot \vec{E} = e \frac{\partial N}{\partial t}$$

VASYLIUNAS' SELF-CONSISTENCY EQUATION

$$\nabla \cdot (\overleftrightarrow{\Sigma} + \overleftrightarrow{\Sigma}_H^*) \cdot \vec{E} = e \frac{\partial N}{\partial t}$$

$$\Sigma_H^* = \frac{eN}{B_I}$$

VASYLIUNAS' STEADY-STATE, 2-D, ZONAL MIC MODEL



**1. THE IMPOSED 2-CELL,
2-D DIPOLE CONVECTION PAT-
TERN DISTORTED BY ANTI-CLOCK-
WISE TWIST IN PLASMAL SHEET
ANNULUS**

**2. FIELD STRENGTH EQUATOR-
WARD OF ANNULUS REDUCED BY
THE ORDER OF Σ_P/Σ_H^* . (THE
SHIELDING EFFECT)**

THE TWO GAMES OF CURRENT EXCHANGE

Both derive their exchange current, $j_{\parallel}$, by taking it from or adding it to $j_{\perp}$ (where $j = j_{\perp} + j_{\parallel} B/B$) according to

$$\nabla \cdot j = 0 \Rightarrow \nabla \cdot j_{\perp} = -\nabla \cdot j_{\parallel} \frac{B}{B} = -B \frac{\partial j_{\parallel}}{\partial s} \frac{B}{B}$$

For the ionosphere, as noted earlier, this is

$$\nabla \cdot (\sum \vec{E}) = j_{\parallel} \sin \chi$$

For the magnetosphere, the operative equation for $j_{\perp}$ is the force balance condition, which for isotropic pressure is:

$$\nabla p = j_{\perp} \times B$$

Solved for $j_{\perp}$, this is

$$j_{\perp} = \frac{B \times \nabla p}{B^2}$$

Giving for the magnetospheric exchange current the expression

$$B \frac{\partial j_{\parallel}}{\partial s} = -\nabla \cdot \frac{B \times \nabla p}{B^2}$$

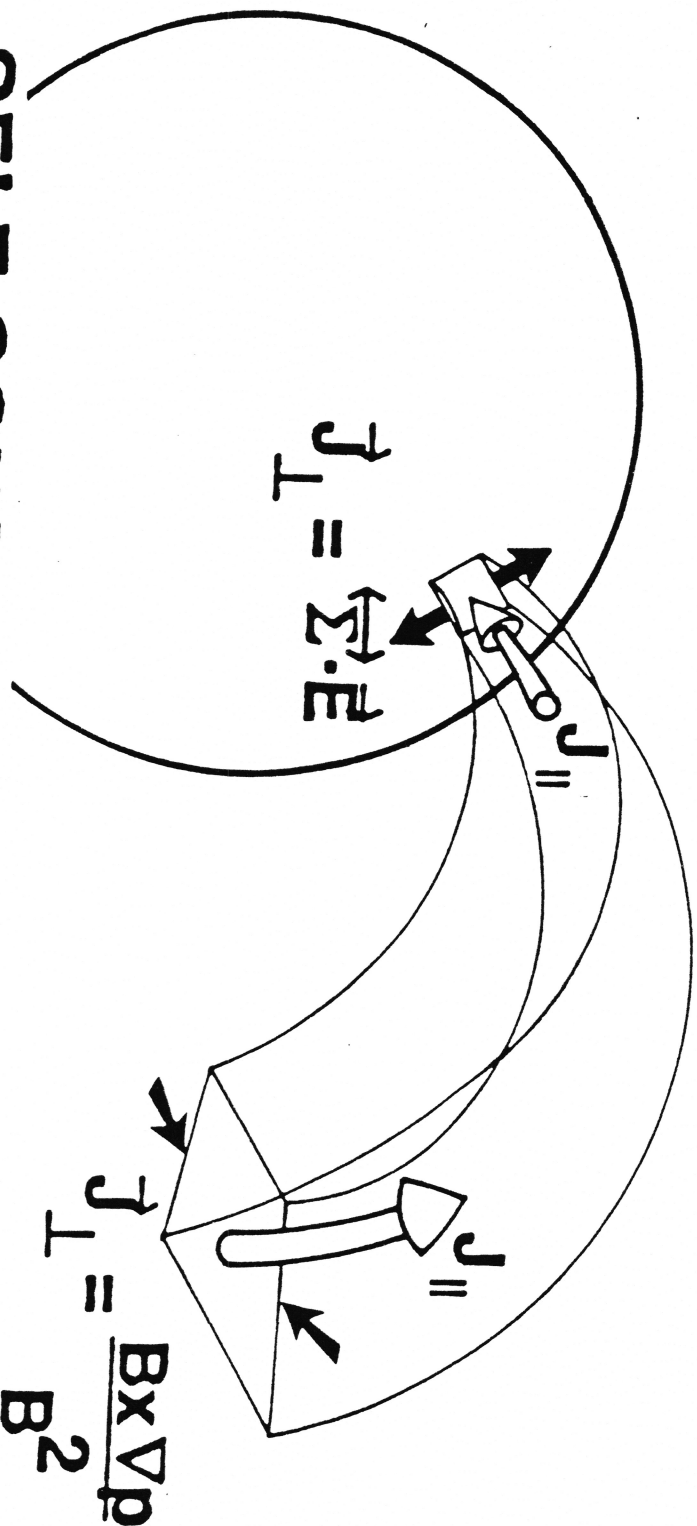
which can be integrated to yield (Wolt, 1983)

$$\text{Magnetospheric game } j_{\parallel} = -\frac{B_{\perp}}{2B_e} \hat{z} \cdot (\nabla_e p \times \nabla_e V)$$

where e and i denote quantities evaluated at the equatorial and ionospheric ends of a flux tube, and

$$V \equiv \int \frac{ds}{B}$$

is the volume of a unit flux tube.

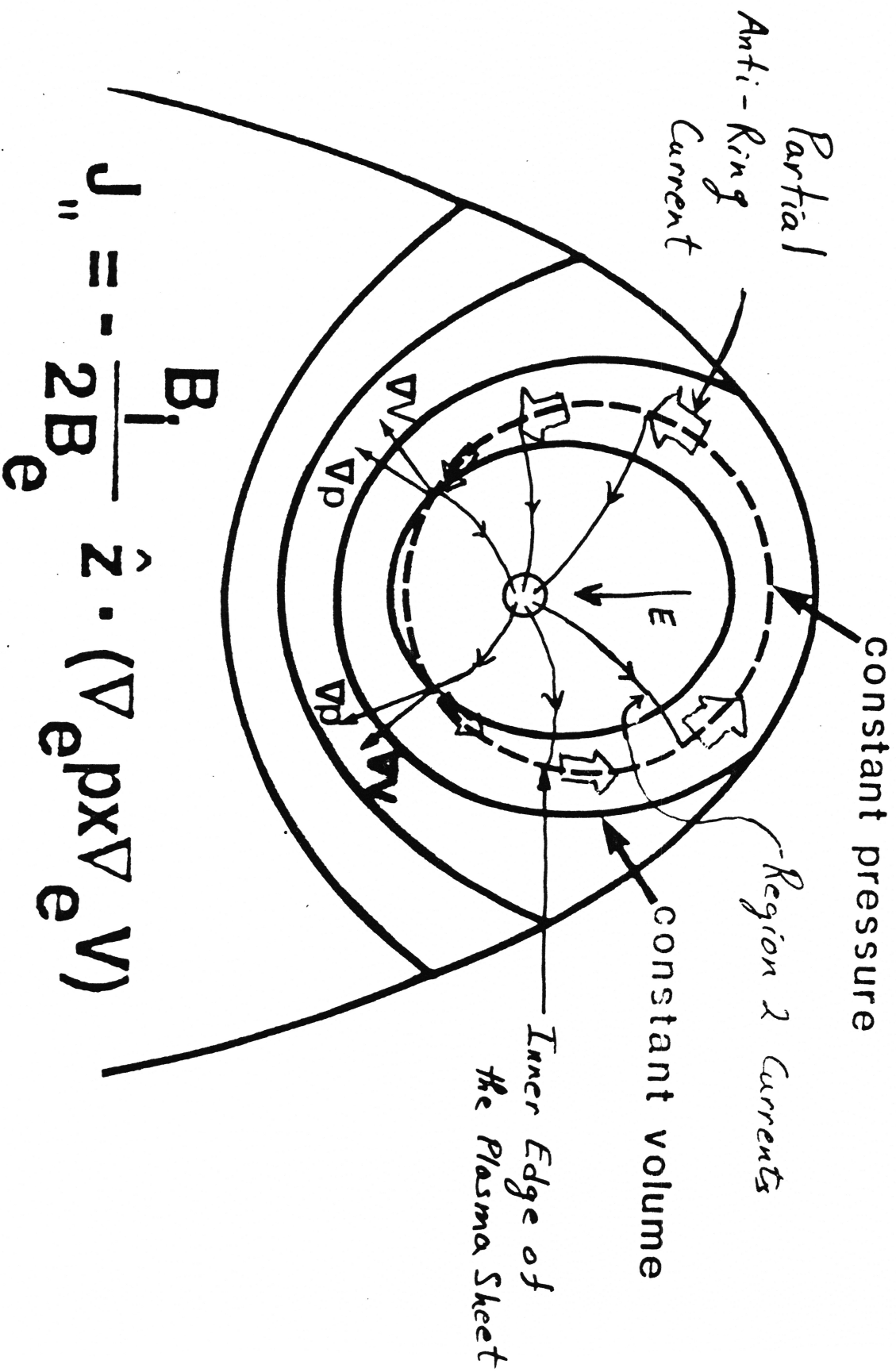


SELF-CONSISTENCY
REQUIREMENT

$$\nabla \cdot \vec{j}_\parallel = -\nabla \cdot \vec{j}_\perp$$

$(j_\parallel)$ IONOSPHERE
 $= (j_\parallel)$ MAGNETOSPHERE

Electric Field Referred
to displace pressure contours



$$\mathbf{J} = -\frac{B}{2B_e} \hat{z} \cdot (\nabla p \times \nabla V)$$

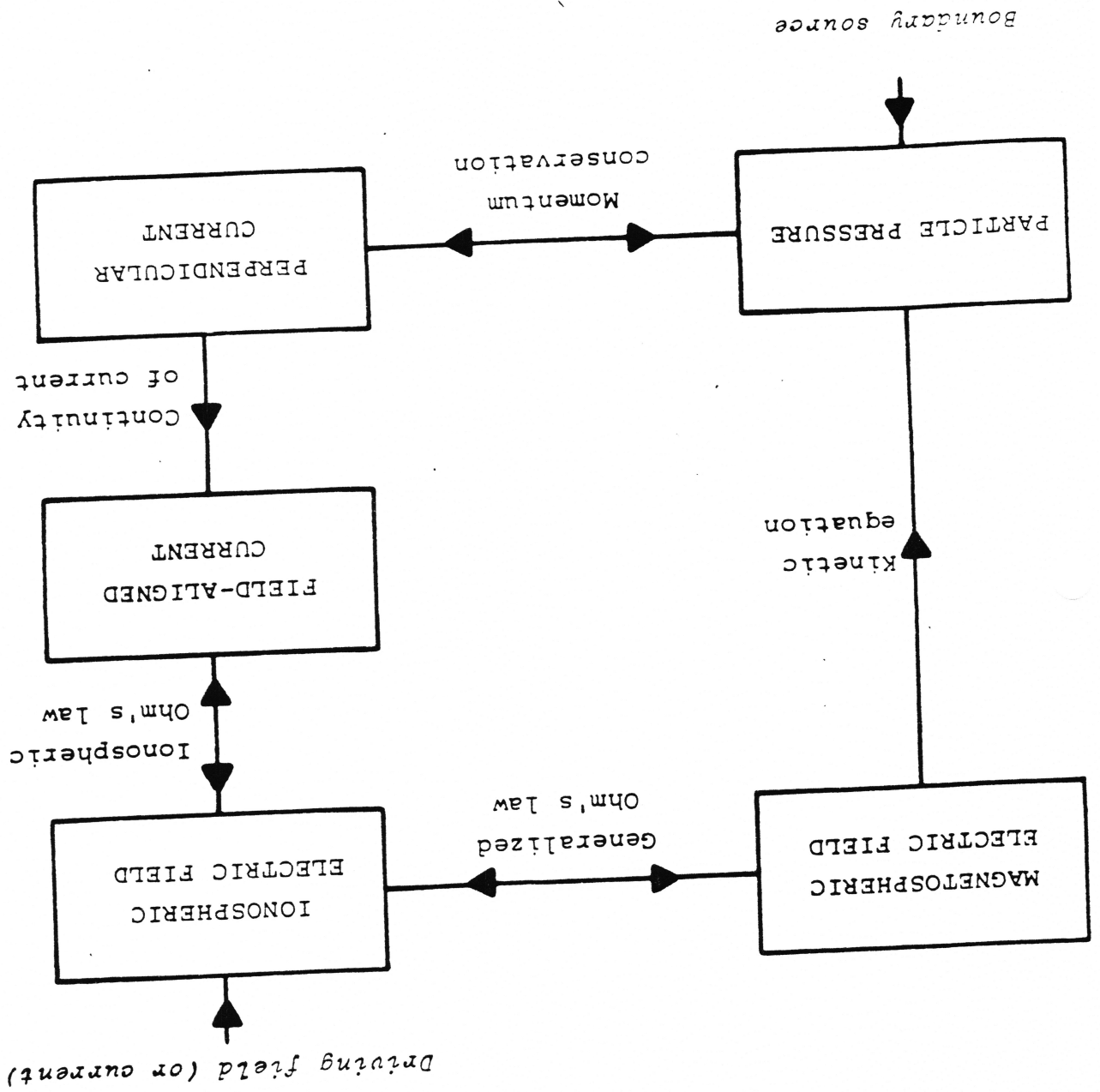
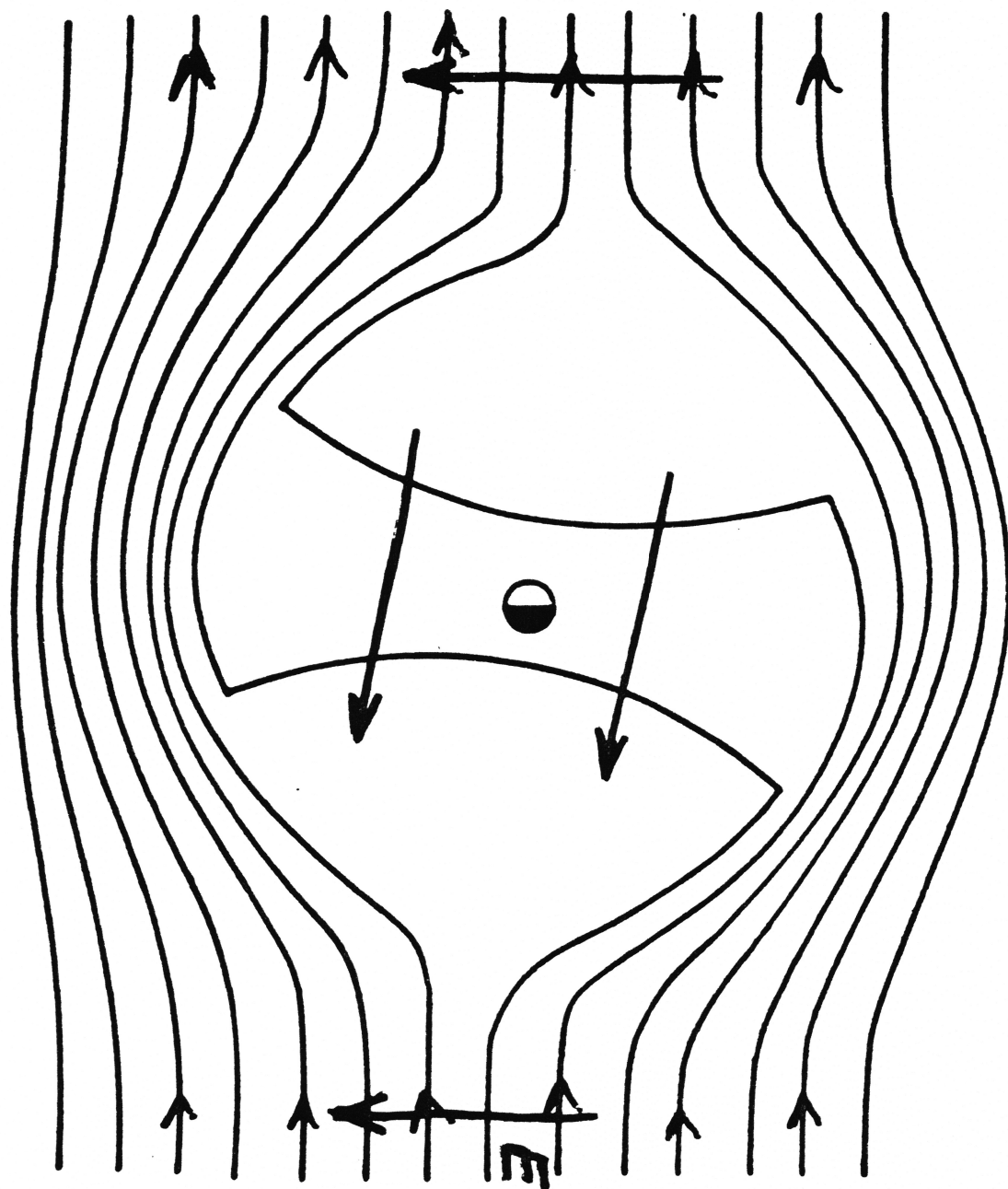
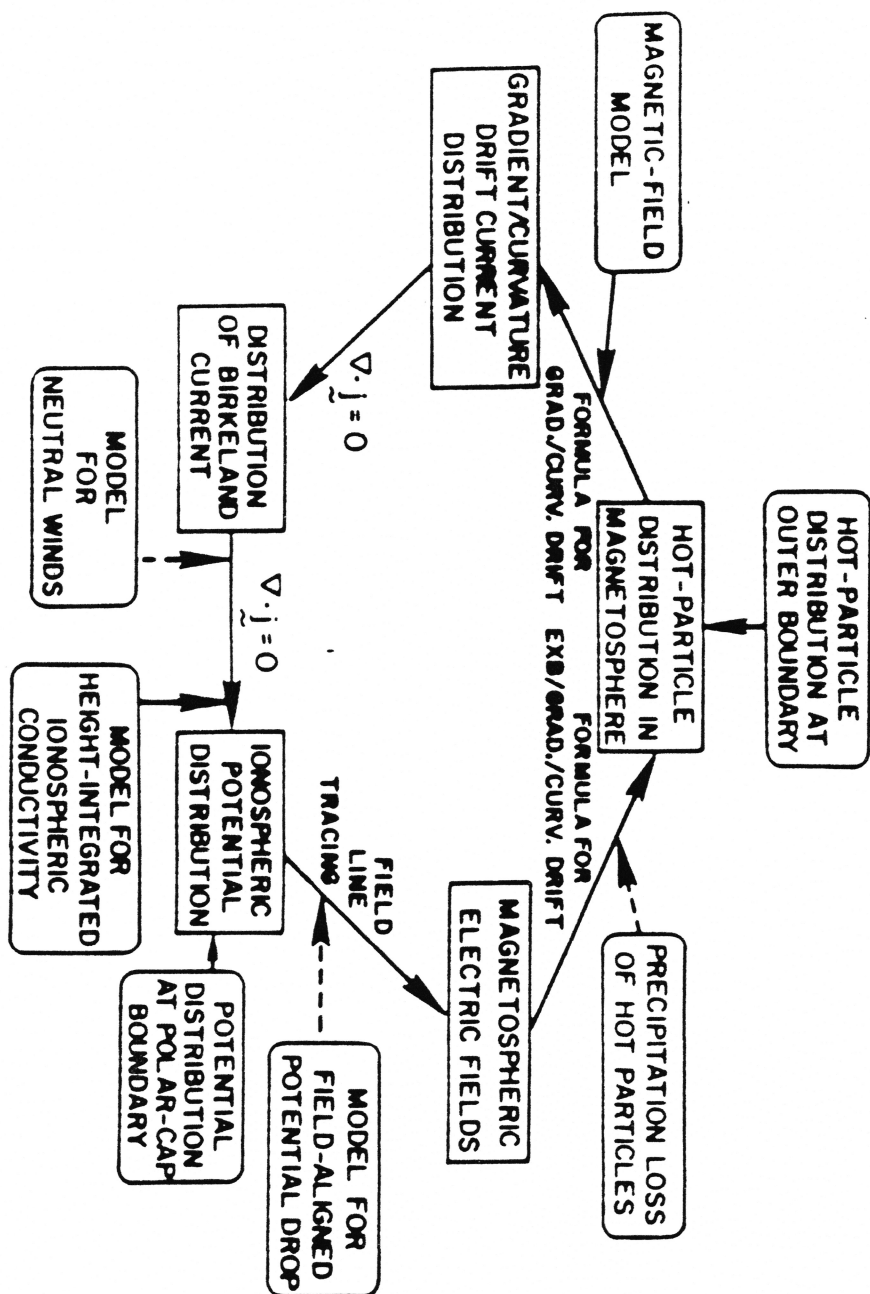
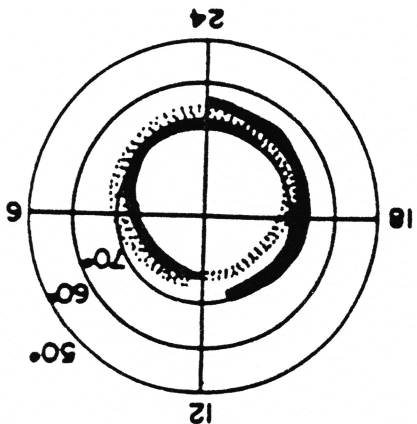


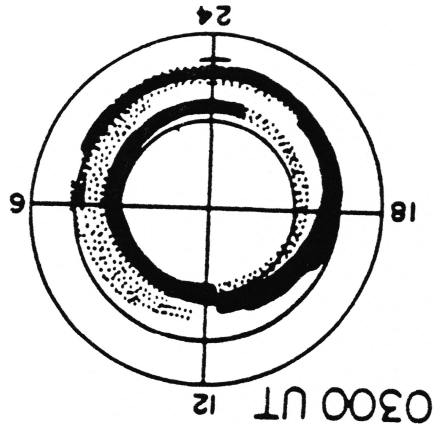
Figure 1





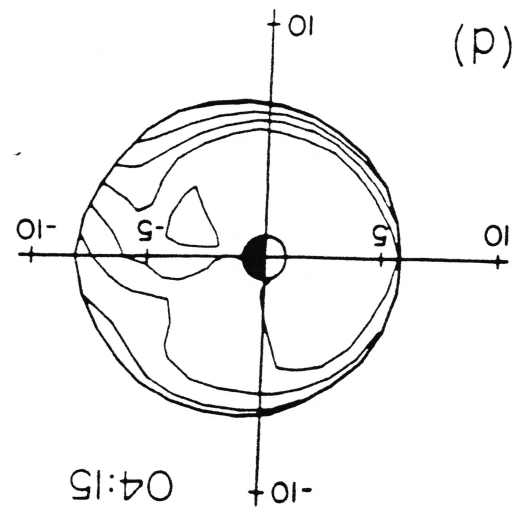
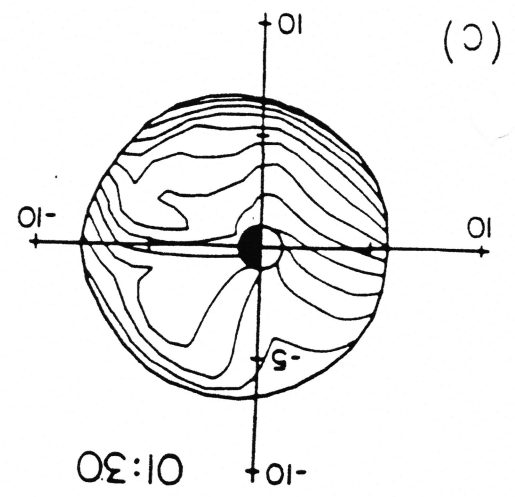
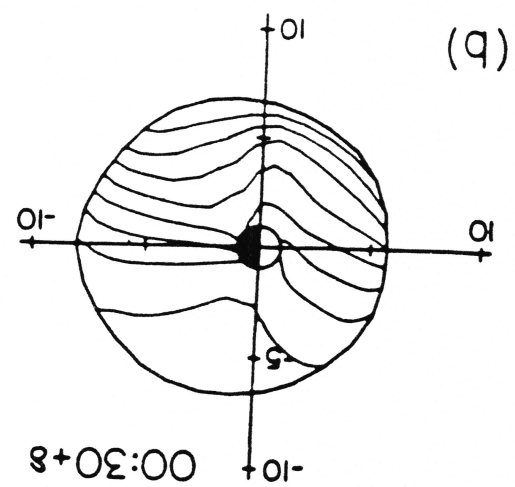
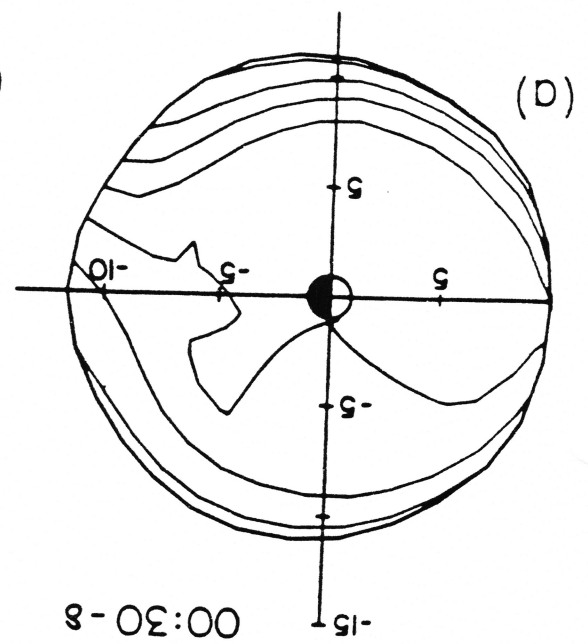


DATA SUMMARY
(IJIJIMA + POTEIRA)

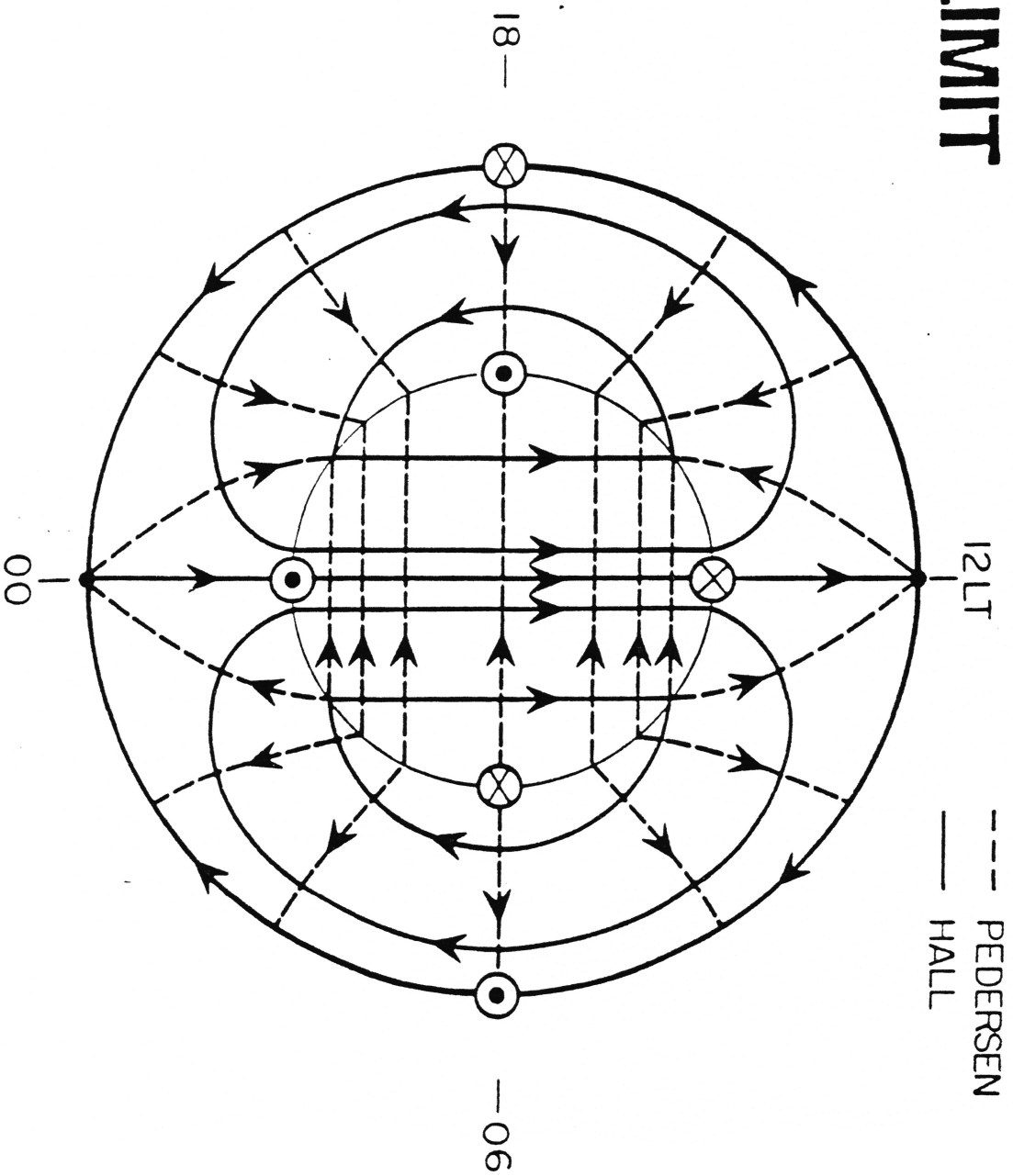


THEORY
(WOLF et al.)

DOWNWARD CURRENT
UPWARD CURRENT



PERFECT SHIELDING LIMIT



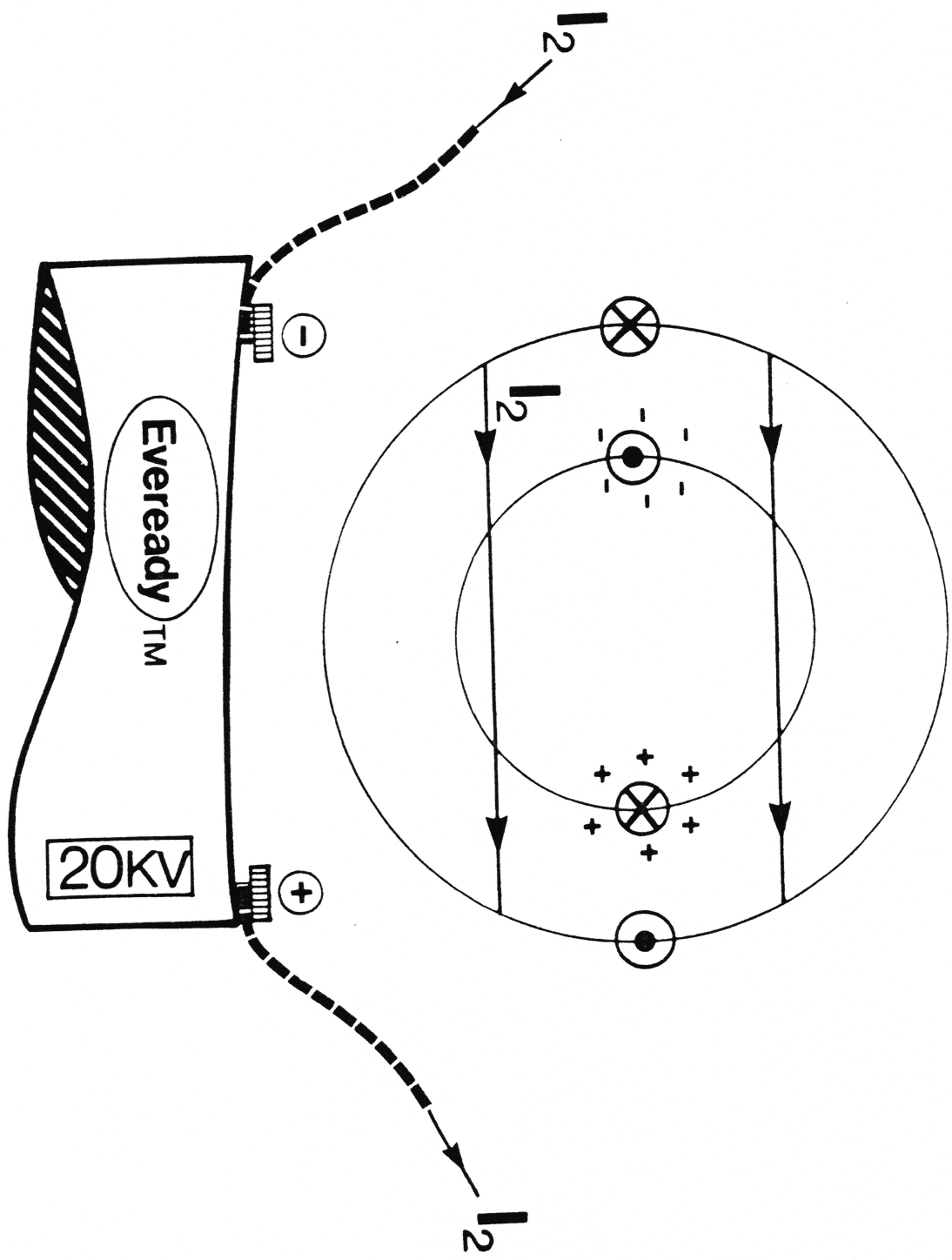
PERFECT SHIELDING LIMIT

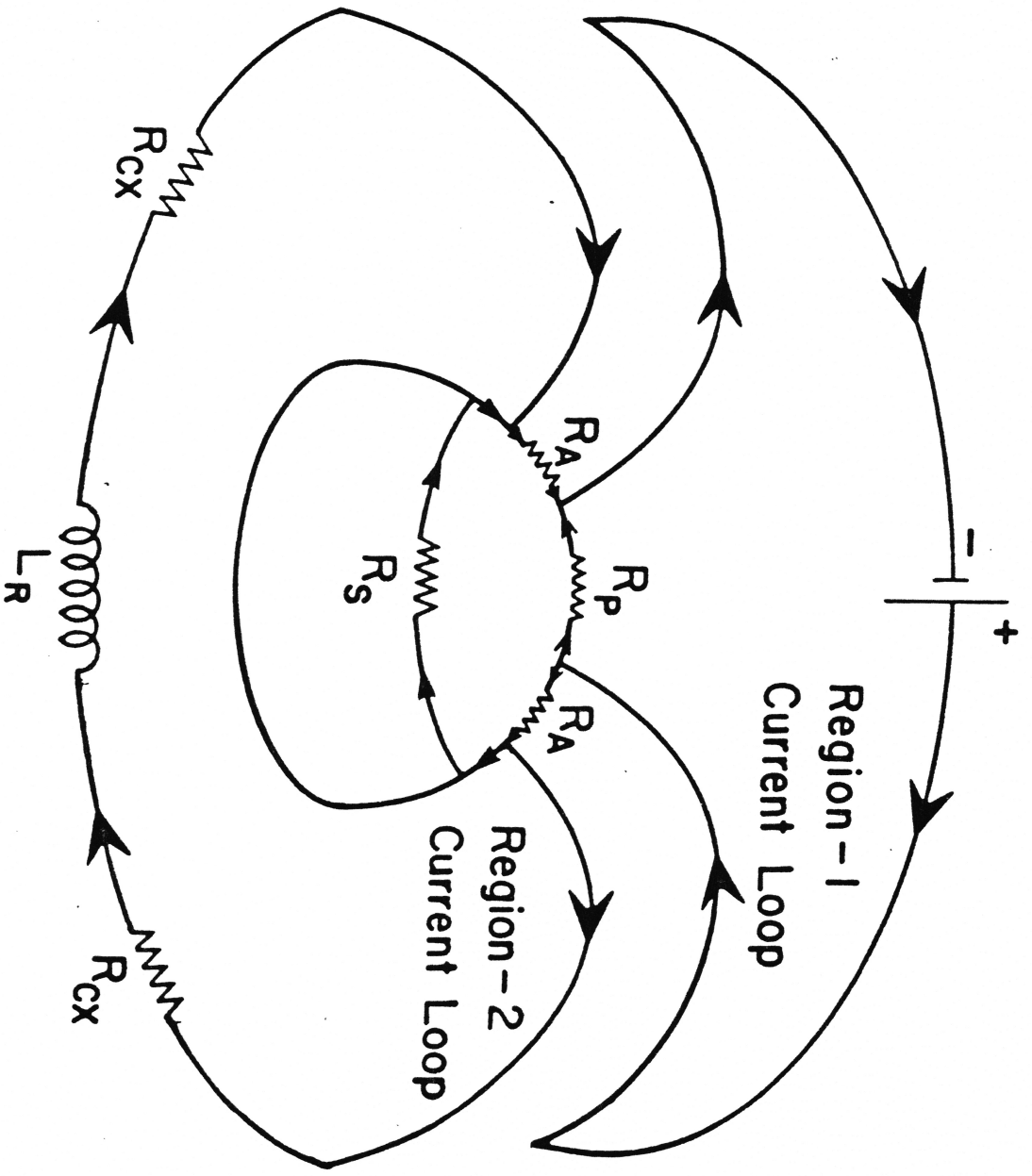
$$I_1^P = \left(\frac{b^2 + a^2}{b^2 - a^2} \right) \Sigma_A^P + \Sigma_e^P \Phi$$

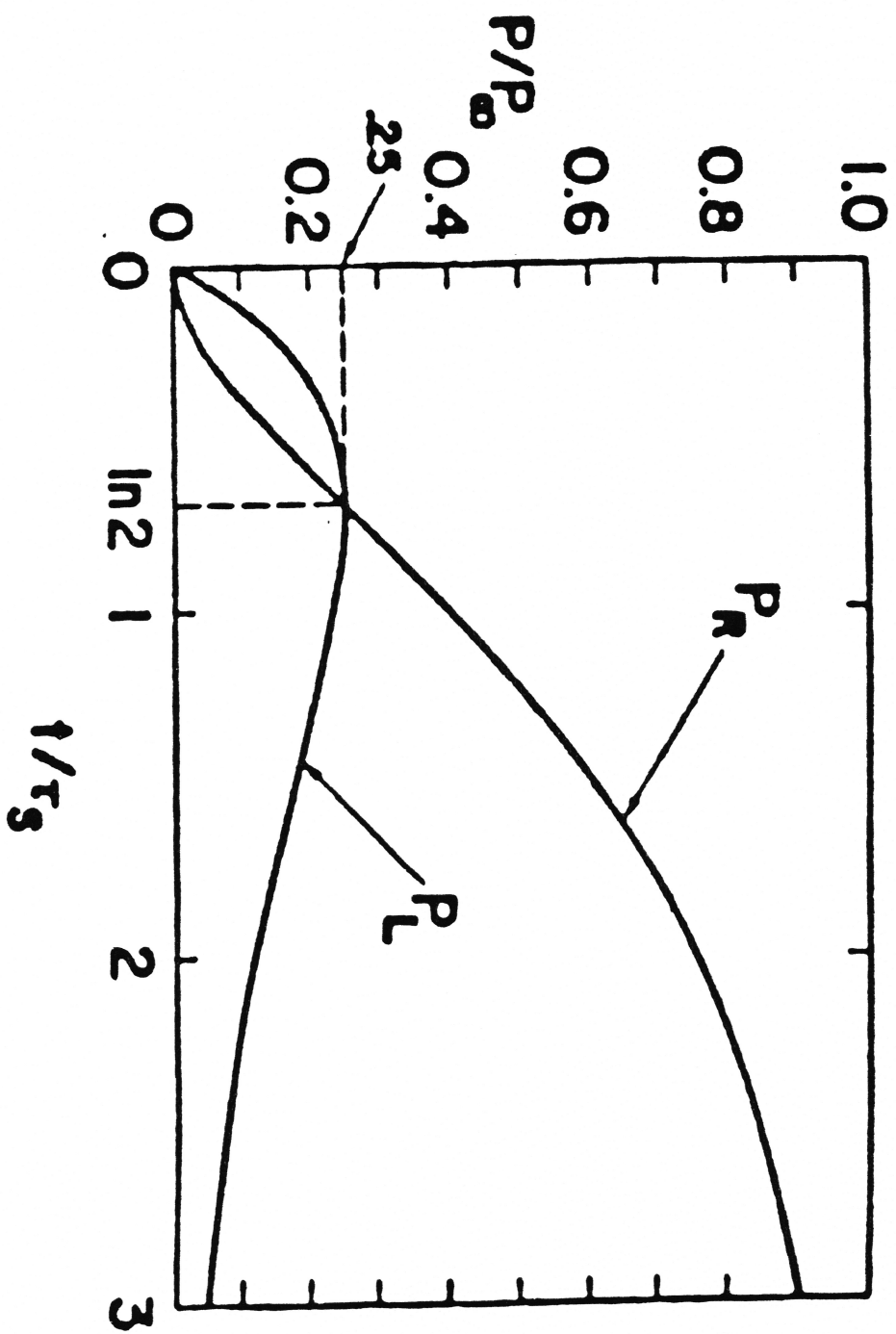
$$I_2^P = \frac{2ab}{b^2 - a^2} \Sigma_A^P \Phi$$

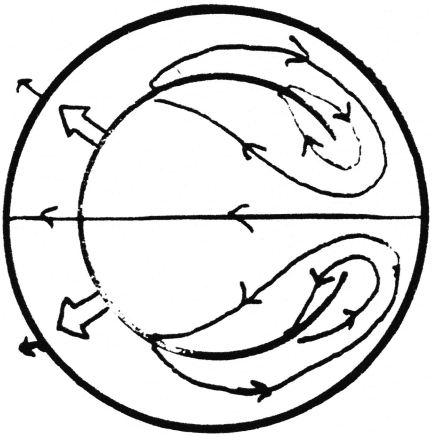
1. BECAUSE OF REGION 2
CURRENTS, THE IONOSPHERE
DRAWS 3 TO 5 TIMES MORE
CURRENT FROM THE REGION
1 DYNAMO

2. I_1 AND I_2 INCREASE
AS THE ANNULAR WIDTH
NARROWS



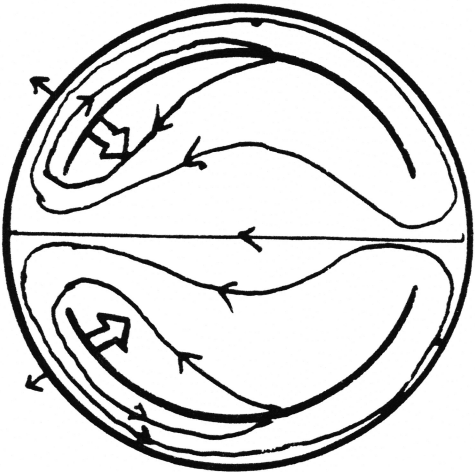






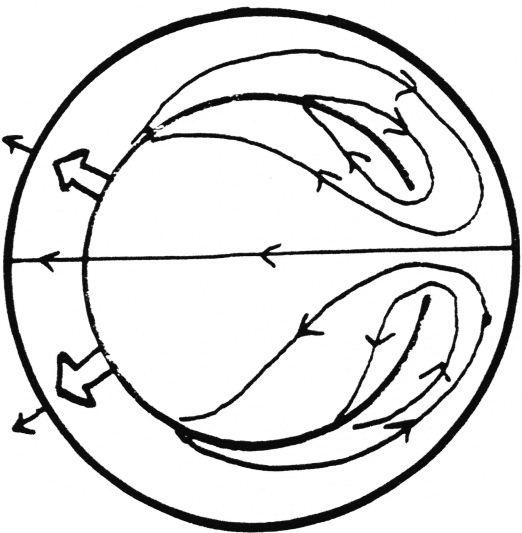
1

Growth
Phase



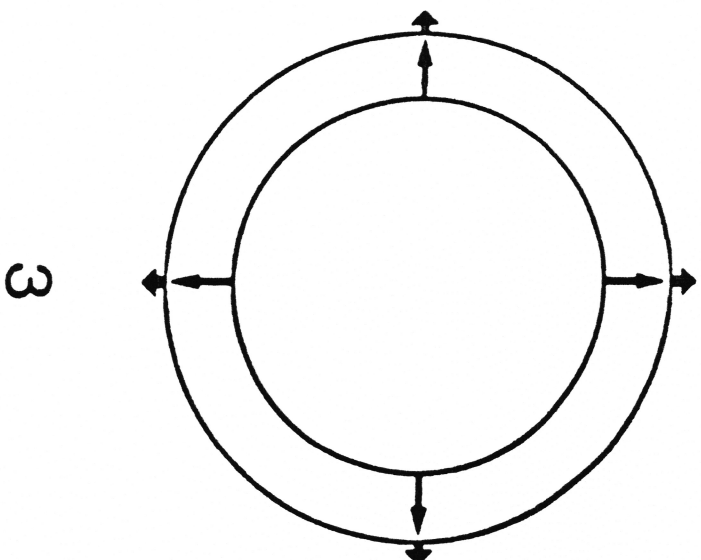
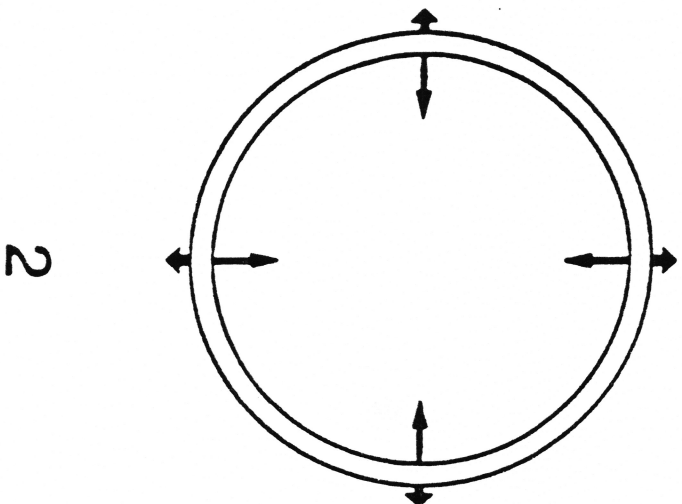
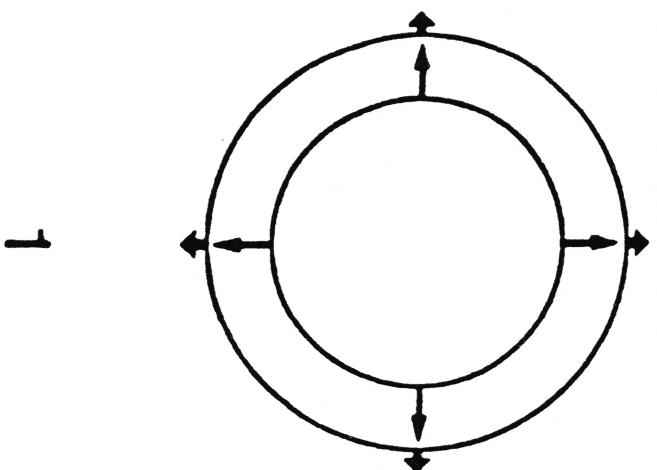
2

Expansion
Phase



3

2nd Growth
Phase



$$\frac{d}{dt}(\pi a^2 B_p) = \Phi$$

COMPARE $\frac{b}{R_e} = (A \frac{a}{b} \Phi)^{3/16}$

WE FIND $\frac{db/dt}{da/dt} \approx \frac{1}{16}$

AND $\frac{b-a}{da/dt} \approx 1 \text{ HOUR}$

FURTHER IMPLICATIONS

- 1. D_{st} DIRECTLY DRIVEN BY Φ**
- 2. SUBSTORMS INDIRECTLY**

FURTHER IMPLICATIONS

$$1. \frac{d}{dt} D_{st} \propto I_2 \frac{d\Phi}{dt} \propto \frac{d}{dt} \Phi^2$$

$$\therefore D_{st} \propto \Phi^2 \propto (IMF B_z)^2$$

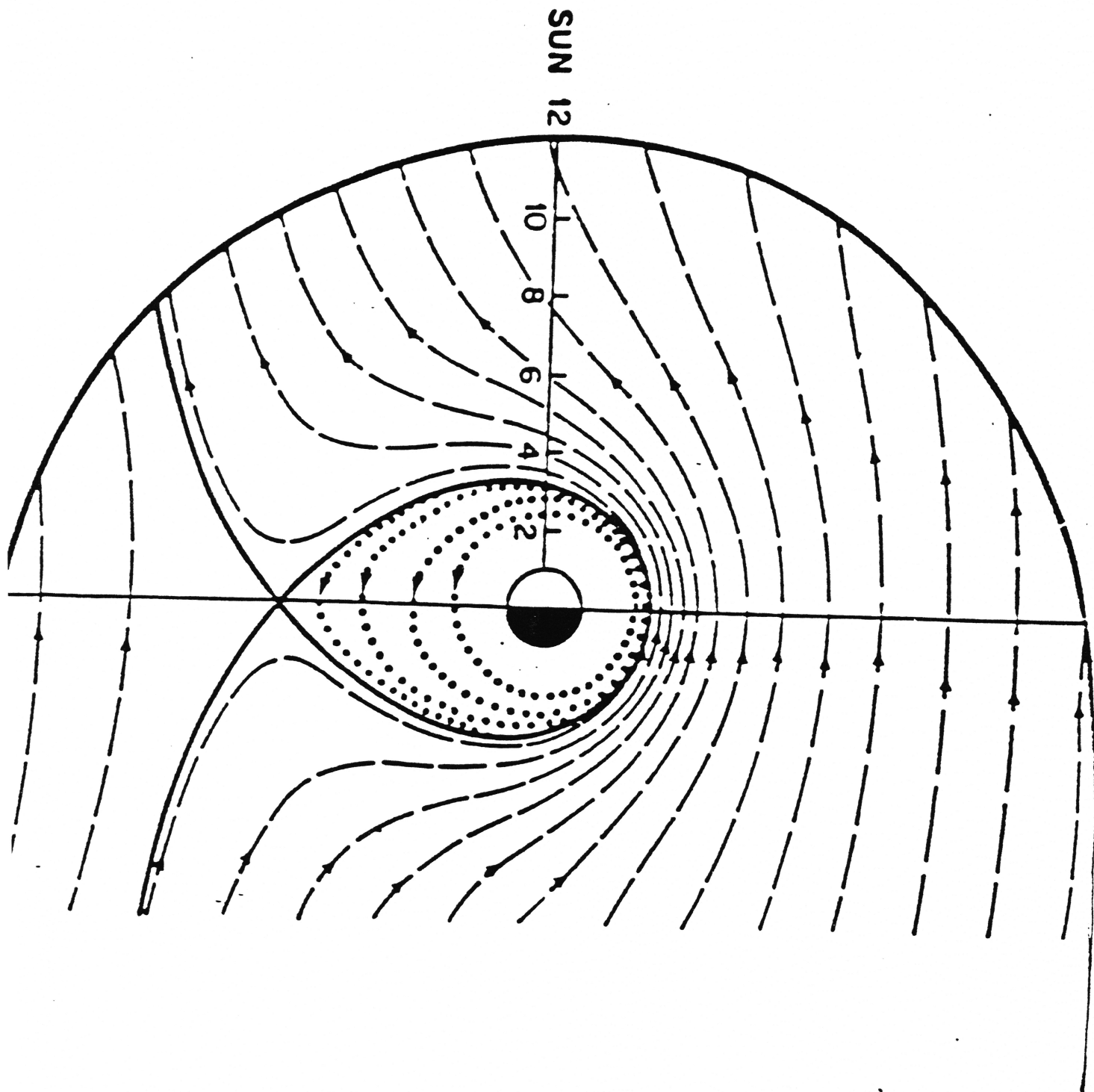
$$2. AE \propto I_1 \propto \Phi \propto IMF B_z$$

CONCLUDING WORDS

The physics of the magnetosphere-ionosphere coupling that takes place on closed, quasi-dipolar field lines is relatively well understood. With the Rice Convection Model (RCM), the subject is well into the computer implementation and application stage. Nonetheless, even here there are still some major and interesting problems. Ionospheric potentials at low latitudes are not what they are supposed to be. Neutral winds stirred by magnetospheric convection may play the dominant role in determining these. Magnetic fields and field aligned currents in the magnetosphere still have to be determined self-consistently. Inductive electric fields can be important and have to be added. But the RCM is now a powerful research tool that can be used to advance global theories and to aid in interpreting global observations.

Lumped circuit analogs of magnetosphere-ionosphere coupling can help understand the system's time dependent global behavior. The property of hot, trapped plasma to simulate an inductor in its interchange of thermal energy eliminates some of the objections to circuit analogs. This avenue should be explored beyond the simple example shown here to see where it goes.

The magnetosphere-ionosphere coupling models are only formulated for the closed field line, quasi-dipolar parts of the magnetosphere. A great but immensely important task is to include the tail. The ultimate task is to attach them to an ionosphere-magnetosphere-solar wind coupling model that spans the polar cap



CURRENT SYSTEM

**REPRESENTATIVE
TOTAL CURRENT (10^6 A)**

Region 1 Birkeland	2.7
Region 2 Birkeland	2.2
Chapman - Ferraro	2.5
Tail	1.5 per $10 R_e$
Cusp	0.2

$$\frac{\partial}{\partial s} \frac{J_u}{B} = 2 \frac{\vec{B}}{3} \cdot (\nabla p \times \nabla B)$$

PRESSURE TERM

$$+ \rho \vec{V} \cdot \nabla \frac{\Omega_u}{B}$$

INERTIAL TERM

